# Vertical Current Shear Retrieval from Shipboard Marine X-band Radar

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#### 1 Data Overview

#### 2 Radar-based Near-surface Vertical Current Shear Measurement Principles

#### 3 Shipboard Marine Radar Near-surface Current Profiling Methodology

4 Results from ITOP Field Campaign



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## Marine X-band Radar System

A science marine radar (MR) system consists of a standard navigation X-band (9.4 GHz) radar that is connected to a PC equipped with a data capture board and analysis software (e.g. WaMoS).

System hardware:



#### Radar antennas:



#### **R/V Roger Revelle:**



## **ITOP Experiment**

Impact of Typhoons on the Pacific (ITOP), western Pacific, 2010. Cruise tracks and mooring locations:



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## **Quasi-Eulerian and Wave-induced Current**

The radar-based current is a superposition of a quasi-Eulerian (wind drift, tidal, geostrophic, and inertial motions) current  $\mathbf{U}_E$  and a wave-induced current  $\mathbf{U}_{Sf}$  (Ardhuin et al. 2009):

$$\mathbf{U}_{R}=\mathbf{U}_{E}+\mathbf{U}_{Sf}.$$

 $U_E$  for deep water linear waves is (Stewart and Joy 1974):

$$\mathbf{U}_E(k_D) = 2k_D \int_0^h \mathbf{U}(z) \exp(-2k_D z) \mathrm{d}z.$$

 $\mathbf{U}_{Sf}$  is given by:

$$\mathbf{U}_{\mathcal{S}f}(k_D,\theta_D) \simeq \mathbf{U}_{\mathcal{S}\mathcal{S}}(f_D) \cdot \mathbf{e}_{\theta_D} + 4\pi k_D \int_{f_D}^{\infty} \int_0^{2\pi} f\cos(\theta - \theta_D) E(f,\theta) \mathrm{d}\theta \mathrm{d}f.$$

 $\mathbf{U}_{SS}$  is the Stokes drift vector for waves with frequencies up to  $f_D$ :

$$\mathbf{U}_{SS}(f_D) = 4\pi \int_0^{f_D} \int_0^{2\pi} f\mathbf{k}(f) E(f,\theta) \mathrm{d}f \mathrm{d}\theta.$$



## **Effective Depth**

Depth weighting function for  $U_E$ :



 $U_E$  represents a weighted-mean of the upper ocean currents.

Linear current profile: Effective depth is 7.8 % of ocean wave length

Logarithmic profile: 4.4 %



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# [Wave and] Current Analysis

Processing steps:

- Backscatter ramp correction
- Pulse-by-pulse georeferencing and trilinear interpolation
- Standard near-surface current retrieval
- Heading correction and current "calibration"
- From 3D spectral density to SNR
- Current profiling

Radar image example:



Goal: Reduce backscatter dependency on range and azimuth.



Corresponding Fourier-fitted ramp:



#### Goal: Reduce backscatter dependency on range and azimuth.





Ramp-subtracted return [×10<sup>2</sup>]

## **Georeferencing and Trilinear Interpolation**

Polar radar images are transformed from ship to geographical coordinates and trilinearly (i.e. in space and time) interpolated. MR image spiral: After trilinear interpolation:



# **3D Dispersion Shell**

Wave energy location in 3D wavenumber-frequency space ("dispersion shell"); still water  $(\mathbf{U} = \langle 0, 0 \rangle \text{ ms}^{-1})$ :

Linear dispersion relationship:

$$\boldsymbol{\omega} = \sqrt{gk} \tanh kh + \mathbf{k} \cdot \mathbf{U}$$

- $\omega$  Angular frequency
- g Acceleration due to gravity
- k Wavenumber
- h Water depth
- k Wavenumber vector
- U Current vector



## **Doppler-shifted 3D Dispersion Shell**

Wave energy location in 3D wavenumber-frequency space ("dispersion shell"); with current  $(\mathbf{U} = \langle -1, 0 \rangle \text{ ms}^{-1})$ :

Linear dispersion relationship:

$$\boldsymbol{\omega} = \sqrt{gk} \tanh kh + \mathbf{k} \cdot \mathbf{U}$$

- $\omega$  Angular frequency
- g Acceleration due to gravity
- k Wavenumber
- h Water depth
- k Wavenumber vector
- U Current vector



Aliasing occurs if a signal is temporally undersampled ( $\omega > \omega_{Ny}$ ). Higher harmonics appear mainly due to nonlinearities in the imaging mechanism (Senet et al. 2001).

Dispersion relationship with higher harmonics:

$$S_{\rho}^{\pm} = \pm (\rho + 1) \sqrt{\frac{gk}{\rho+1}} \tanh\left(\frac{kh}{\rho+1}\right) + \mathbf{k} \cdot \mathbf{U}$$

Aliasing symmetry conditions:  $P(k_x, k_y, \omega) = P(k_x, k_y, \omega + n\omega_{Ny})$  $P(k_x, k_y, \omega) = P(-k_x, -k_y, -\omega)$ 

- *S* Angular frequency
- *p* Integer for harmonic order
- P 3D image power spectrum
- *n* Integer for frequency intervals



### **Spectral Signal of Near-surface Current**

Current fit minimizes distance of 3D dispersion shell from wave signal (Young et al. 1985; Senet et al. 2001).

Frequency slice through 3D spectrum (1.12 rads<sup>-1</sup>):



Wavenumber-frequency slice from 3D spectrum ( $k_x = 0$  radm<sup>-1</sup> and 0 radm<sup>-1</sup>  $\leq k_y \leq k_{Ny}$ ):



#### Ship-track Dependent Current Errors I

Shipboard current measurements require accurate heading data. Errors induce a spurious cross-track current:  $U_{\perp} = U_s \sin \theta_e$  (Pollard and Read 1989).



Solution: (1) Correct gyro heading using multi-antenna GPS (King and Cooper 1993), and (2) perform water-track "calibration" to determine (constant) radar–compass misalignment angle (Joyce 1989).

Ship maneuver offering calibration opportunity:

Alignment error: tan  $\alpha = \frac{\langle \delta u'_{d} \delta v_{s} - \delta v'_{d} \delta u_{s} \rangle}{\langle \delta u'_{d} \delta u_{s} + \delta v'_{d} \delta v_{s} \rangle}$ 



Us	Vs	δus	$\delta v_s$
-0.164	-4.354	-0.089	-2.181
0.015	0.008	0.089	2.181

u' <sub>d</sub>	$v'_d$	δu' <sub>d</sub>	$\delta v'_d$
-0.135	3.662	-0.001	2.13
-0.133	-0.597	0.001	-2.13



#### From 3D Spectral Density to Signal-to-Noise Ratio I

Background noise as  $f(k, \omega)$  is estimated from spectral half with lowest standard deviation of spectral density over *k*.



Goal: Identify wave signal over broad range of directions and wavenumbers.





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Time series of MR (0.225 radm<sup>-1</sup> or ~2 m) and ADCP (21 m) currents.



Scatter plots and statistics for MR-ADCP comparison.

Current speed:



## **Near-surface Current Profile Examples I**



Example of MR near-surface vertical current shear measurements with WW3-based filtered Stokes drift.



#### **Near-surface Current Profile Examples II**

"Parallel flow" and "clockwise forcing".



#### **Near-surface Current Profile Examples III**

"Counter flow" and "counter-clockwise forcing".



#### **Near-surface Current Profile Examples IV**

"Parallel flow" transitioning f/ "counter-" to "clockwise forcing"; flow reversal.



## **Choice of Background Current Depth I**



Time series of differences between MR (0.11 radm<sup>-1</sup> or ~4.5 m) and ADCP (21 m) currents.

## **Choice of Background Current Depth II**

Current MR (0.11 radm<sup>-1</sup>) - ADCP (101 m) Nind RR1010 **RR1012 RR1014** RR1015 100 Speed [cms<sup>-1</sup> / ms<sup>-1</sup> 80 60 4٢ 20 300 Direction [°] 200 100 15 20 25 0 5 10 Time [days]

Time series of differences between MR (0.11 radm<sup>-1</sup> or ~4.5 m) and ADCP (101 m) currents.

## **Stokes and Ekman Drift**

Mean background-current-corrected MR profiles.



## **Conclusions & Outlook**

- Compared MR near-surface currents with ADCP reference measurements
- MR currents are in good agreement with reference
- Presented first MR near-surface current profiles
- Near-surface current profiles' response to wind and wave forcing agrees with our physical expectations (flow to the right of the wind, speed decays, deflection angle increases with depth)
- Future work / outlook:
  - Determine exact effective depths of near-surface current profiling results through numerical inversion of U<sub>E</sub>(k<sub>D</sub>)
  - Validate MR near-surface current profiles, e.g. using drifters
  - Apply methodology to further MR data sets to study Stokes drift and Ekman dynamics



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## **Spectral Sensitivity to Currents**

Technique's sensitivity, in terms of wavenumber resolution cells:



## **Spectral Sensitivity to Bathymetry**

Sensitivity to bathymetry:



### **MR Near-surface Current Profile Accuracy**

Mean deviations of MR currents from 5th degree polynomial fit:



#### Mean SNR Across Dispersion Curve

Mean SNR along lines covering a range of radial distances from the dispersion curve, for different frequencies:



## **Marine Radar Current Shear Measurements**







# Currents, Winds, and Waves



Difference between ASHTECH

Conventional mechanical gyro compasses are reliable but have errors of O(1 °) depending on ship speed, heading, and latitude (Bowditch 2002).



Gyro compass error dependency on heading:

# **ITOP Experiment**

Impact of Typhoons on the Pacific (ITOP), western Pacific, 2010.

Air-Sea Interaction Spar (ASIS) buoy (Graber et al. 2000):



Extreme Air-Sea Interaction (EASI) buoy (Drennan et al. ASIS-EASI buoy pair: 2014):

**R/V Roger Revelle with** 





Photo credit: Hans C. Graber ASIS / EASI graph credit: Henry Potter



## Stratification

Wind-driven currents are heavily affected by stratification (Price et al. 1987; Price and Sundermeyer 1999).



#### Stokes Drift Estimated from WAVEWATCH III

Scatter plots and statistics for EASI-N– and EASI-S–WW3 non-directional Stokes drift speed comparison (Ardhuin et al. 2009).



WAVEWATCH III peak wave parameters, 12 August 2010, 12:00 UTC. Significant wave height, currents (black Peak wave period and direction: arrows) and winds (white):







## **JCOPE-T Modeling Example**

JCOPE-T ocean model – a regional tides resolving version of POM (Miyazawa 2012) – output for vertical (color code) and horizontal current (arrows) at 2 m, 12 August 2010, 12:00 UTC.



## **ASIS-S Near-surface Current Measurements I**

Time series of ADV (4 m) and ADCP (21.22 m) measurements:





Scatter plots and statistics for ADV-ADCP comparison.



#### Spectral Analysis of ASIS-S ADV Data

Rotary ADV spectra (Gonella 1972); the inertial period is 35.54 h:

